

SEISMIC VELOCITIES AND SUBSEA PERMAFROST IN THE BEAUFORT SEA, ALASKA

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The distribution of high-velocity material was used as an indicator of ice-bonded permafrost. Observations from ice survey and marine seismic records, coupled with control from a small number of drill holes, suggest that ice-bonded permafrost is extremely widespread in the Beaufort Sea. Large areas of high-velocity material at shallow depths, 10-40 m below the seabed, were observed near Prudhoe and Harrison Bays. In some cases these zones extended up to 35 km from shore. It was also common to find that depths to the high-velocity material increased with distance from the shore. Observed depths were as great as 150-200 m below the seabed. Velocities also commonly decreased with distance from shore to a point where it was no longer possible to identify ice-bonded horizons. Reflection analysis revealed several deep near-horizontal reflectors that re-occurred along segments of the coastline from Harrison Bay to the Canadian border. Reflectors at approximately 200 m and 450 m could be permafrost-related, although no control exists for an evaluation of this interpretation. Locally these reflectors may extend out to the shelf edge to approximately the 100-m water depth.

Petroleum industry seismic data were used to obtain an improved understanding of the regional distribution of subsea permafrost. Our observations in the U.S. Beaufort Sea are based on data from Harrison Bay east to the Canadian border. The most detailed information is from the coast out to a water depth of approximately 15 m, with the greatest detail from the Prudhoe and Harrison Bay regions (Figure 1).

It has been demonstrated that refraction analysis of the upper part of deep seismic records can be useful in locating high-velocity zones that may be ice-bonded permafrost (Munter et al. 1976, 1978). However, not all frozen materials can be located. Due to their low velocities, which are caused by warm subsea permafrost temperatures, saline pore water, gas in voids, and the coarse resolution of this technique.

Reflection interpretations can also be used to acquire information on permafrost distribution and structure, but the lack of velocity data usually makes these interpretations more speculative than those from refraction analysis. Beyond the 15-m water depth, out to the shelf edge, our only information is from reflection analysis.

METHODS

The early-return data used for this study included nonproprietary records from Western Geophysical Company, Geophysical Service Inc., as well as data released by British Petroleum. Records from the land-sea transition were acquired during the winter over the sea ice. Examples of marine and ice survey records showing high-velocity zones interpreted as ice-bonded permafrost can be found in Neave and Sellmann (1982) and Munter et al. (1976).

Control for these observations is very limited and it is virtually nonexistent beyond the 15-m water depth and in shallower water outside the Harrison and Prudhoe Bay regions. The existing control is from shallow drilling, penetrometer observations, and high-resolution seismic studies (Osterkamp and Harrison 1976, 1980, 1981; Harrison and Osterkamp 1981; Blouin et al. 1979; Sellmann and Chamberlain 1979; Miller and Bruggers 1980; Rogers and Morack 1980, 1981). Observations in the Canadian Beaufort Sea provide indirect control with a range of permafrost velocities, horizontal limits, and thicknesses (Munter et al. 1976).

Error estimates were made to verify the analytical methods and measurements (Neave and Sellmann 1982). The consistency of results from duplicate coverage, often acquired during different years and seasons, also helped develop confidence in the results.

A model and appropriate equations are required to interpret the records and obtain depth and velocity data for the construction of maps and sections. The model used was a simple plane homogeneous layer over a half space, and our equations are derived from those provided by Grant and West (1965). Minor modifications to these equations are reported in Neave and Sellmann (1982).

A simplifying assumption was used for the refraction depth determinations: the water layer was combined with the low-velocity bottom sediments to make a single upper layer. An upper-layer velocity of 1.8 km/s was used for all profiles out to the 15-m water depth. In this zone, the upper-layer velocities were observed to range from 1.6 to 2 km/s. In deeper water an upper-layer velocity of 1.5 km/s was used.

The spatial resolution of data from this study is obviously not as great as could be obtained from

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